Use of Zircon U-Pb Ages And Bulk Trace Element and Nd-Hf Isotope Compositions to Determine Soil Provenance in a Limestone Terrane, Middle TN, USA

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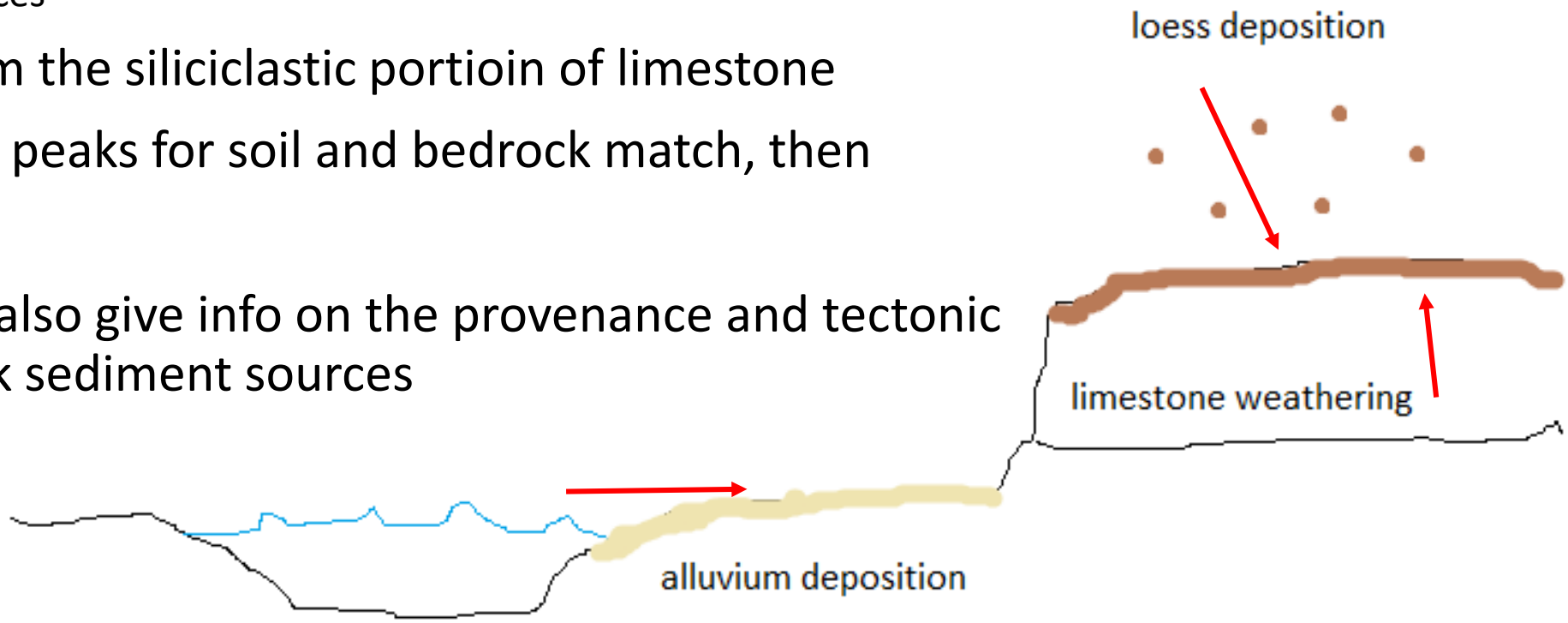
How did soil form in middle TN?

- In-situ chemical weathering of limestone?
- Input of exotic material such as alluvium or loess?
- Identify the source (provenance) of soil-forming material by studying samples of
  - undisturbed soil (high, flat, non-agricultural areas)
  - potential soil sources
- Obtain zircon from the siliciclastic portioin of limestone
- If zircon U-Pb age peaks for soil and bedrock match, then

soil formed in-situ

- Zircon U-Pb ages also give info on the provenance and tectonic setting of bedrock sediment sources



Why zircon U-Pb dating?

- Elemental and isotopic composition and mineralogy not reliable provenance indicators
  - Transport of loess and river alluvium can change composition and mineralogy of deposit
  - Chemical weathering during soil formation changes composition and mineralogy
- Zircon age spectra less likely to be affected by transport or chemical weathering
- Only two previous studies of soil provenance using zircon U-Pb dating; the first was:

Brimhall, G.H., Lewis, C.J., Compston, W., Williams, I.S., Reinfrank, R.F., 1993. Darwinian zircons as provenance tracers of dust-size exotic components in laterites: mass balance and SHRIMP ion microprobe results, in Soil Micromorphology: Studies in Management and Genesis. Elsevier, pp. 65–81. http://dx.doi.org/10.1016/S0166-2481(08)70398-2

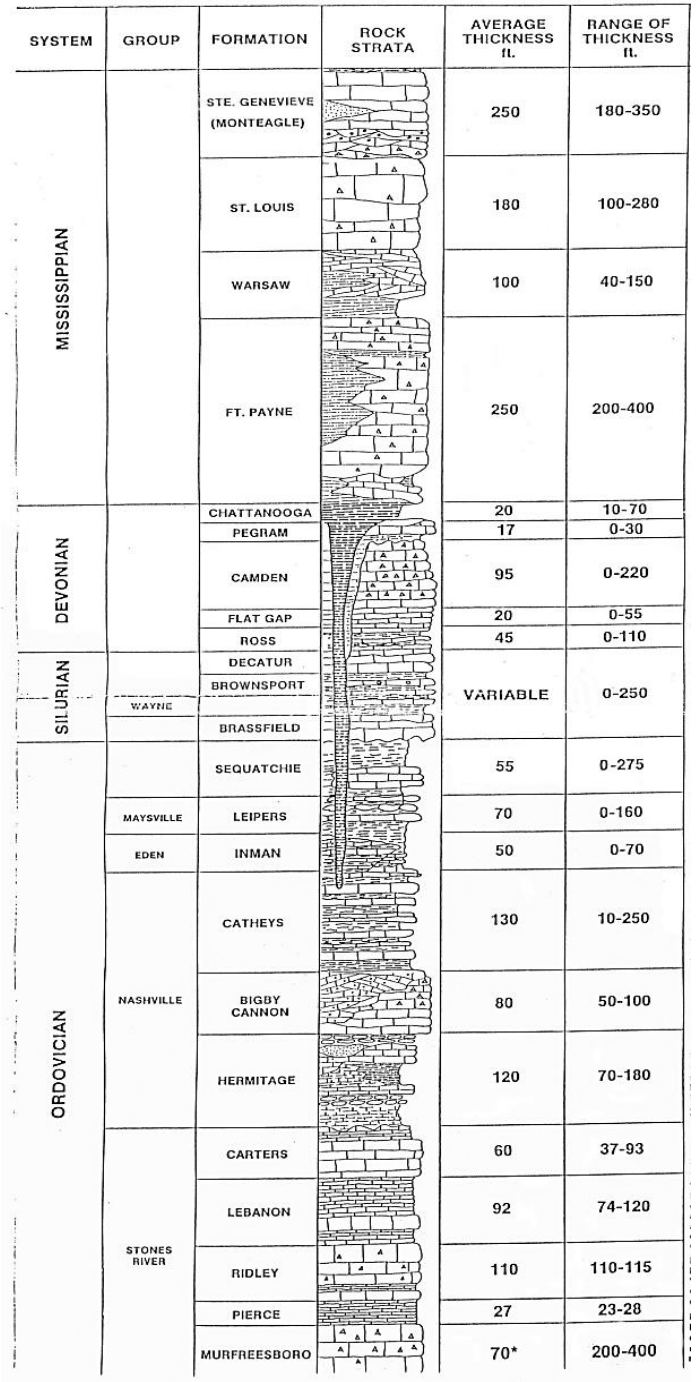
Middle Tennessee Geology

Site 1: Fort Payne Formation

Site 2

Bigby-Cannon Limestone

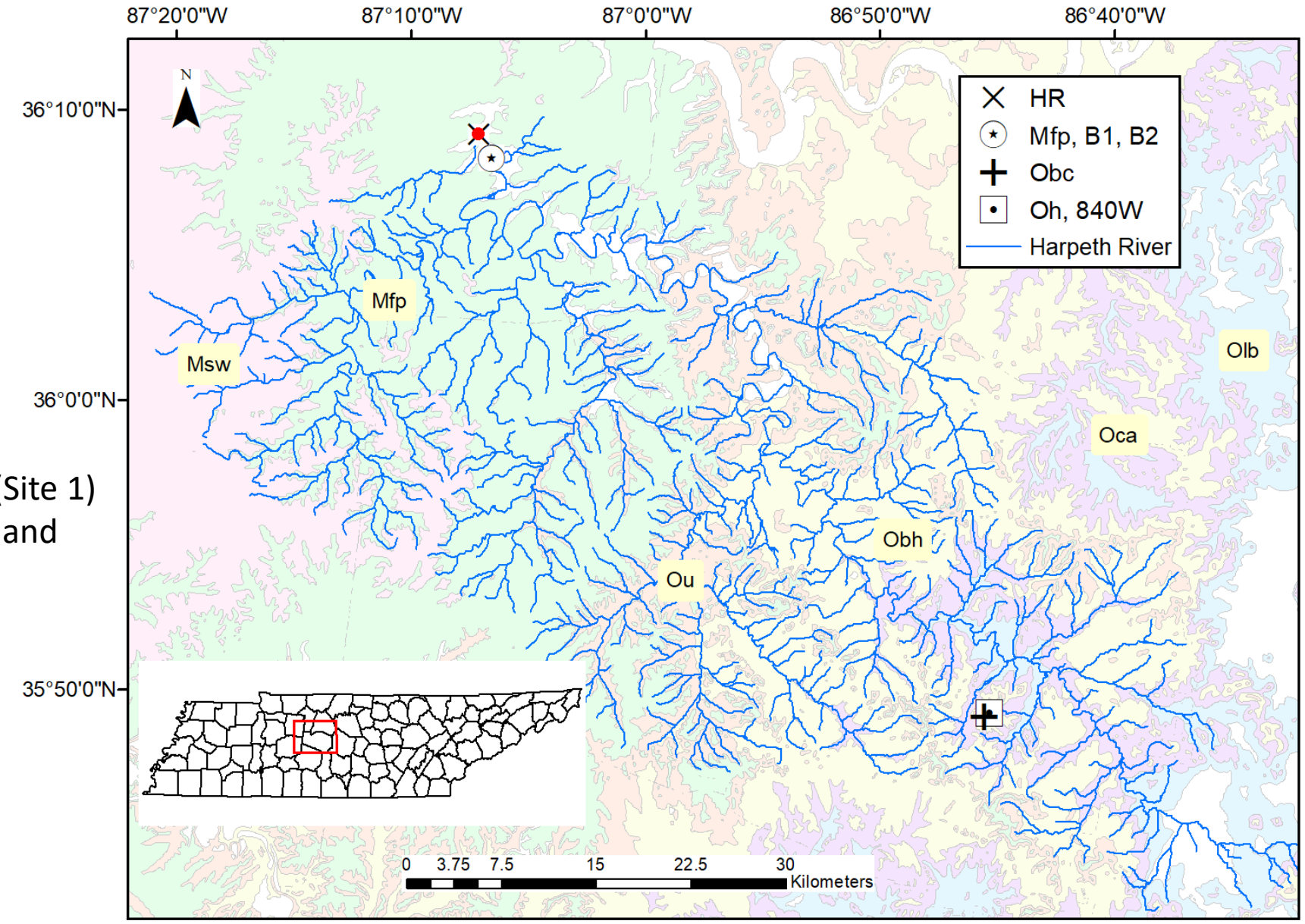
Hermitage Formation



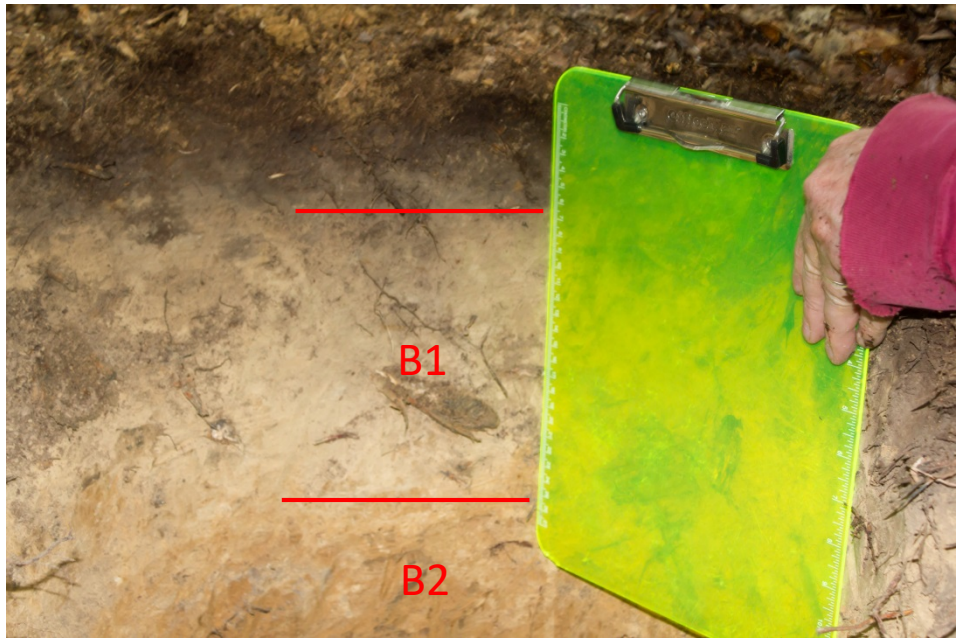
SYSTEM	GROUP	FORMATION	ROCK STRATA	AVERAGE THICKNESS ft.	RANGE OF THICKNESS ft.
MISSISSIPPIAN		STE. GENEVIEVE (MONTEAGLE)		250	180-350
		ST. LOUIS		180	100-280
		WARSAW		100	40-150
		FT. PAYNE		250	200-400
DEVONIAN		CHATTANOOGA		20	10-70
		PEGRAM		17	0-30
		CAMDEN		95	0-220
		FLAT GAP		20	0-55
		ROSS		45	0-110
SILURIAN		DECATUR		VARIABLE	0-250
		BROWNSPORT			
	WAYNE				
		BRASSFIELD			
ORDOVICIAN		SEQUATCHIE		55	0-275
	MAYSVILLE	LEIPERS		70	0-160
	EDEN	INMAN		50	0-70
	NASHVILLE	CATHEYS		130	10-250
		BIGBY CANNON		80	50-100
		HERMITAGE		120	70-180
	STONES RIVER	CARTERS		60	37-93
		LEBANON		92	74-120
		RIDLEY		110	110-115
		PIERCE		27	23-28
		MURFREESBORO		70*	200-400

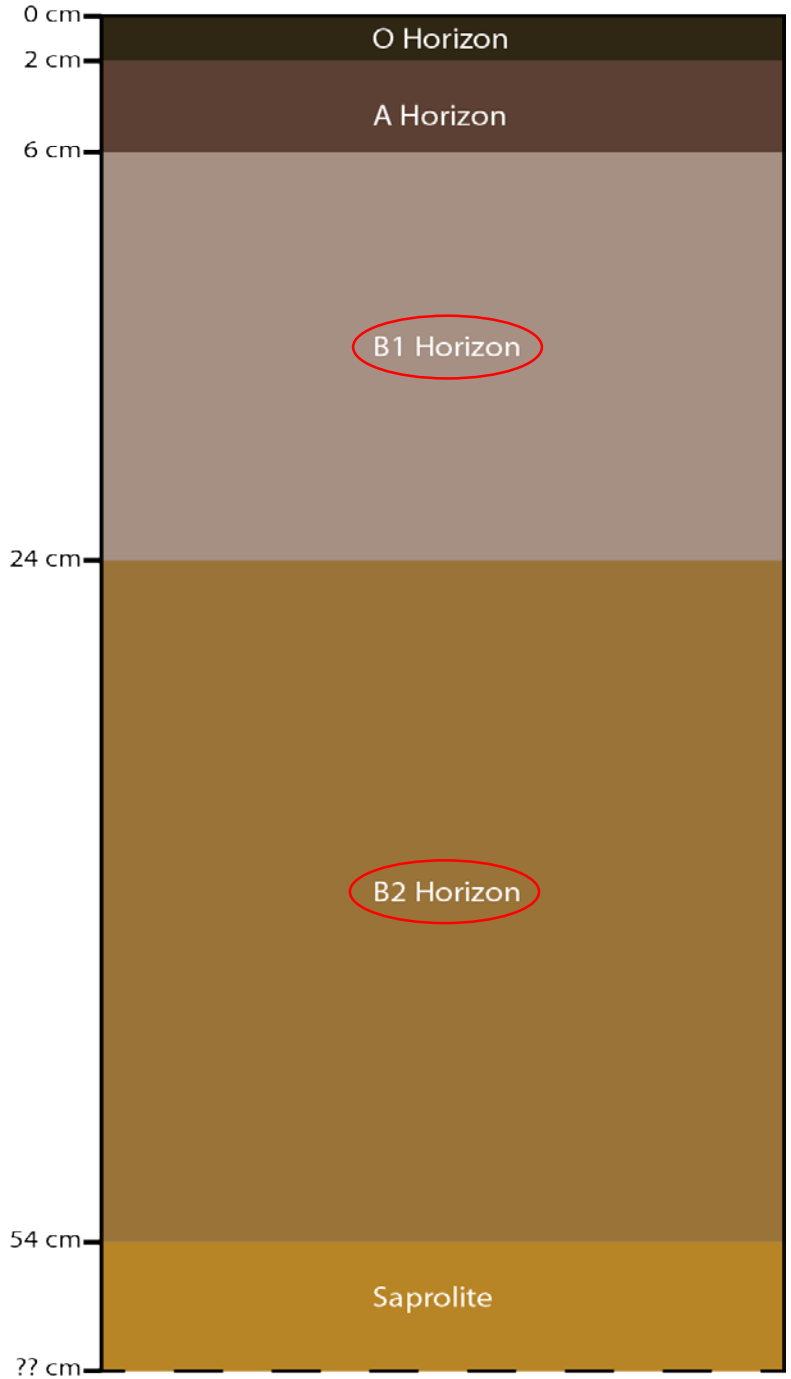
Sample Locations

- Mfp = Mississippian Fort Payne (Site 1)
- Obh = Ordovician Bigby Cannon and Hermitage limestones (site 2)



Field Site 1: Harpeth River Terrace: Ultisol





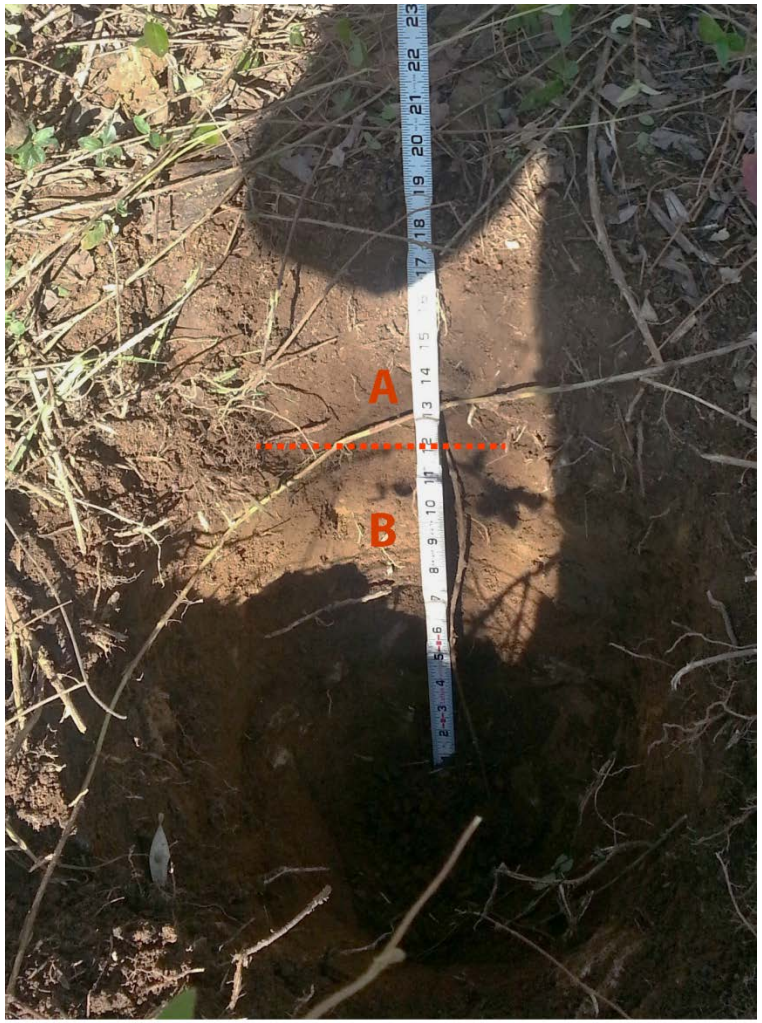
Field Site 2: Highway Outcrop

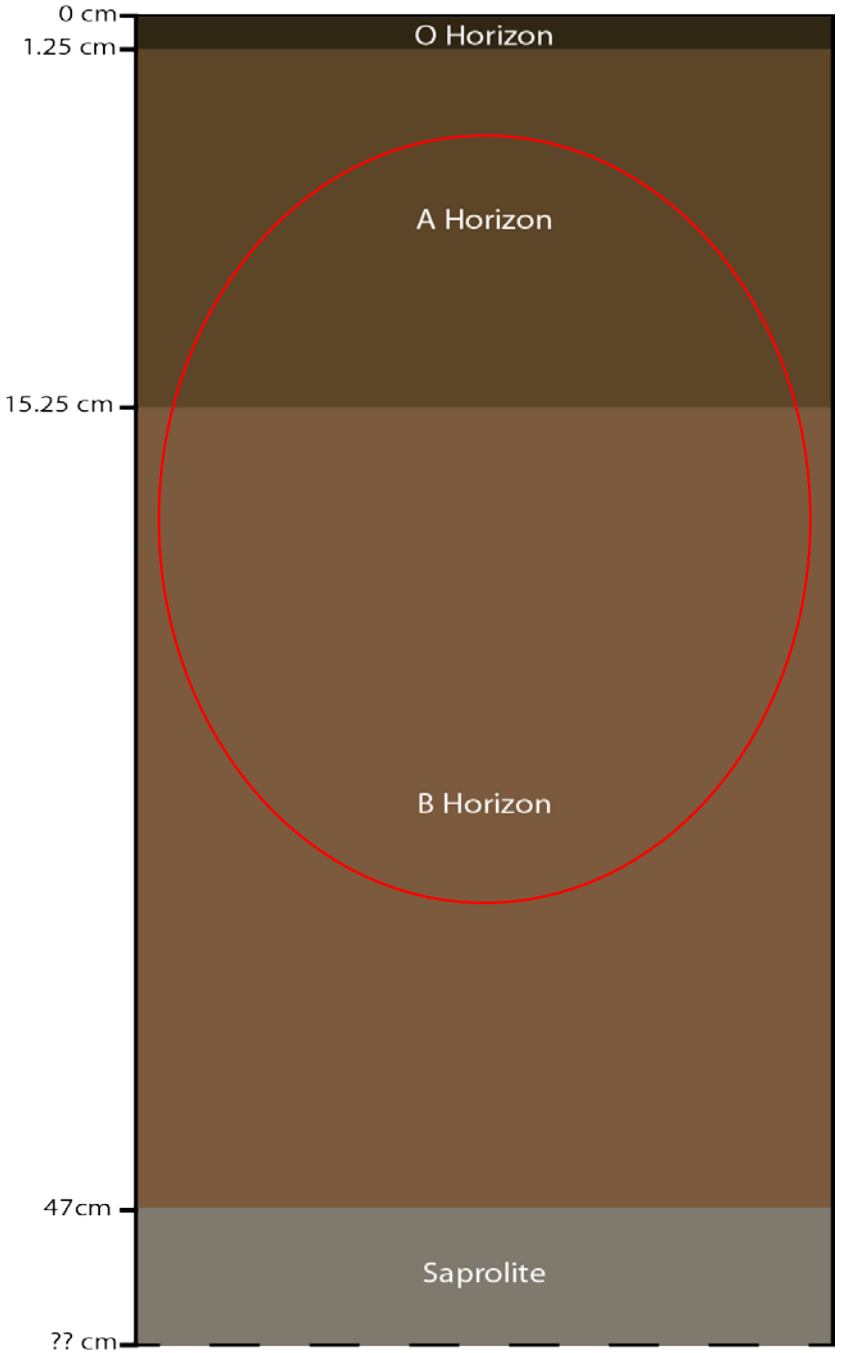
Sampling

- Alfisol soils atop Hermitage Fm. (Oh) Ordovician sandy and argillaceous limestone bedrock

Field Site 2:

Sampling of alfisol





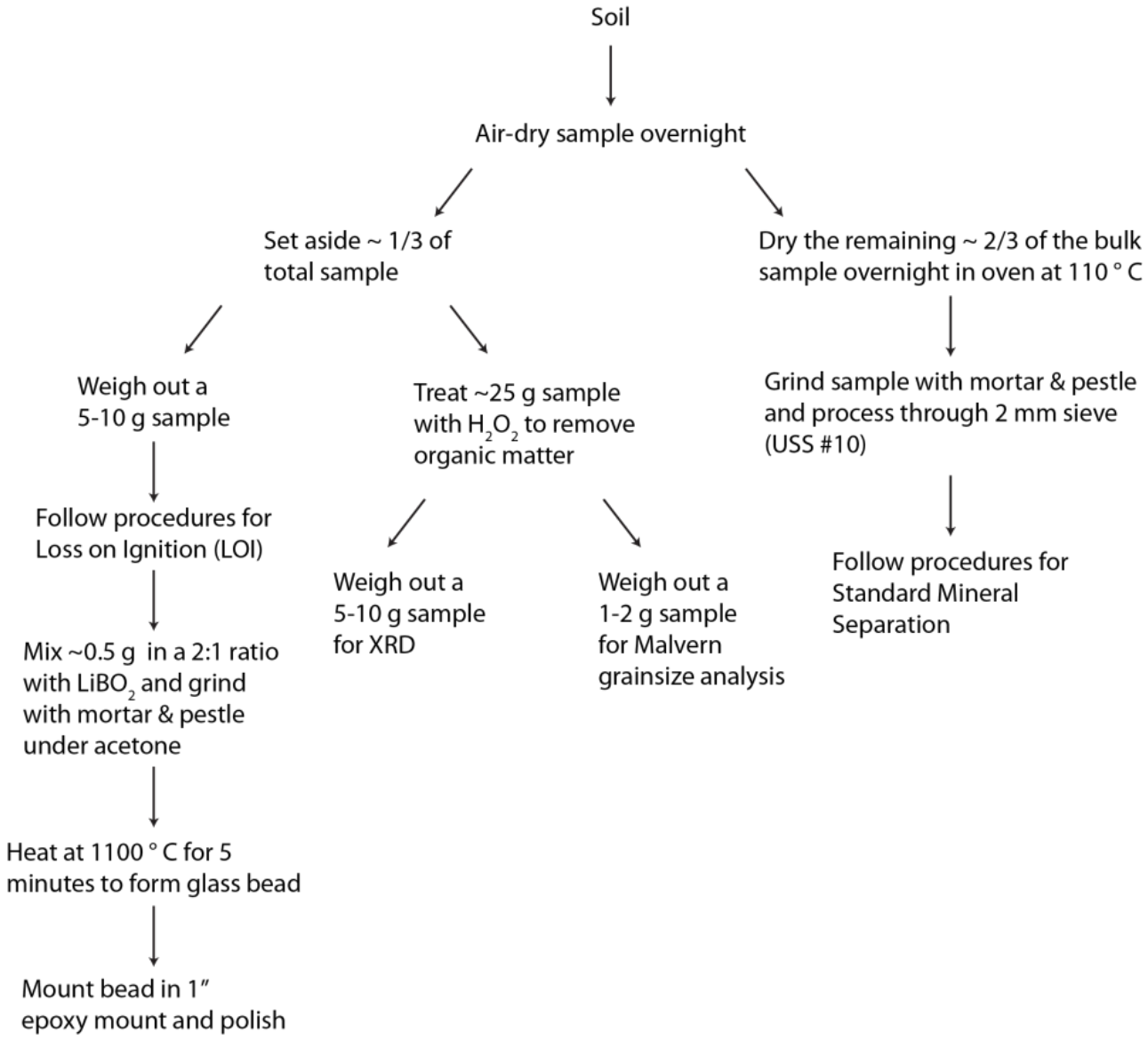
Field Site 2: Obc sample

- Bigby-Cannon Limestone (Obc) Ordovician phosphatic calcarenite/limestone
- Located 0.15km away from the Hermitage-Alfisol sample site

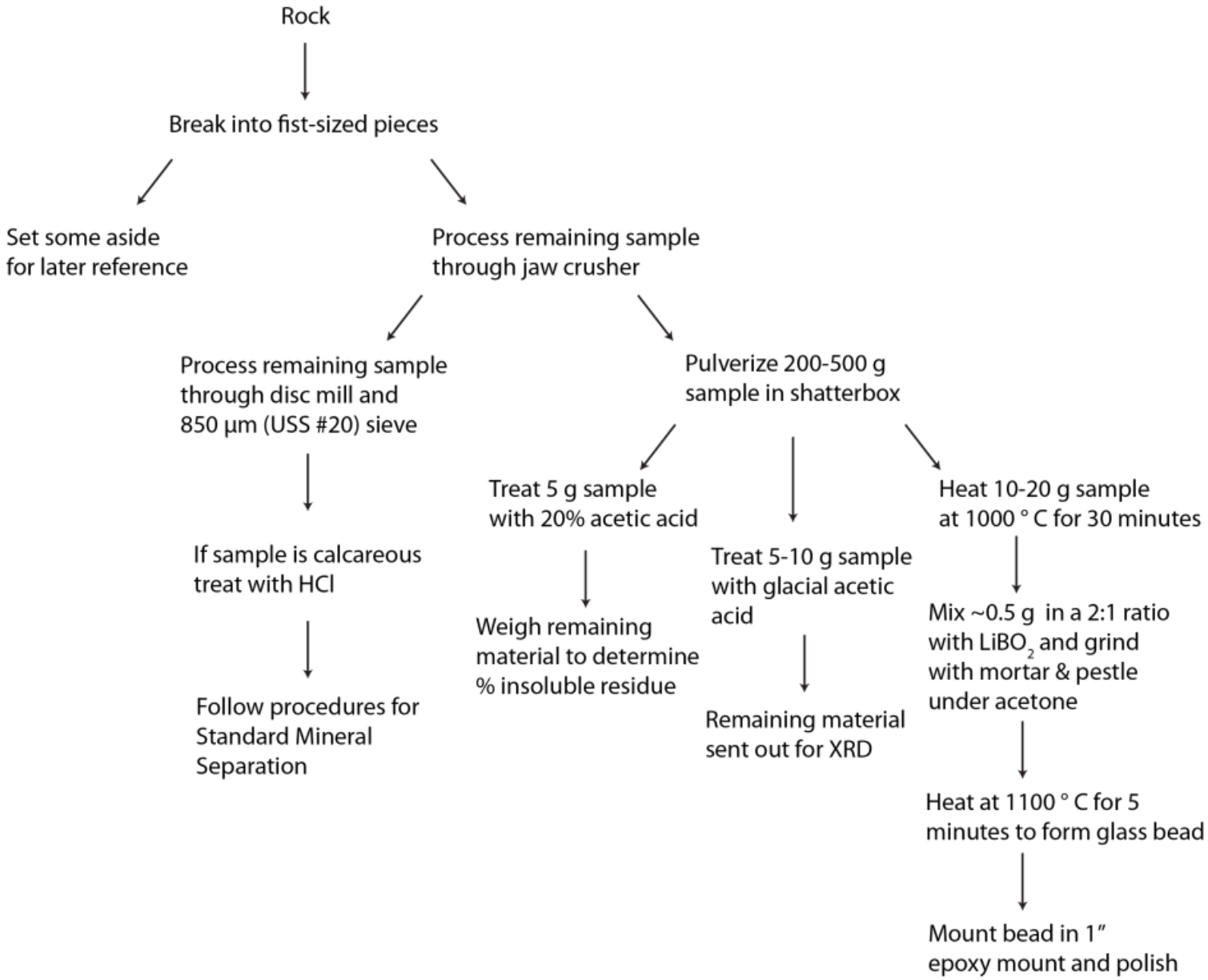
Xiaomei Wang and Jenna Nam

Loess outcrop in Meeman Shelby Park, Memphis, TN

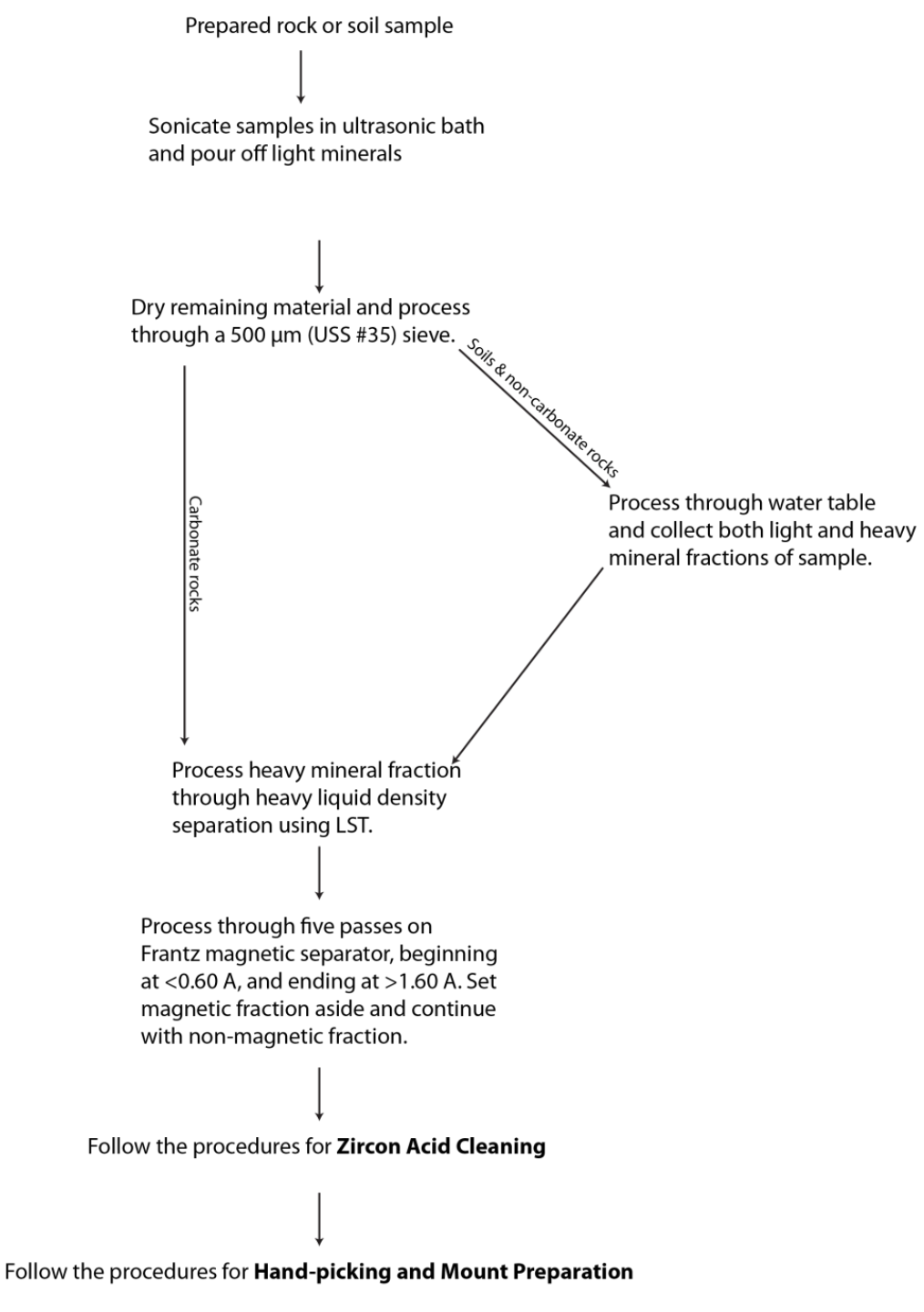
Sample Preparation – Soil & Loess



Sample Preparation - Rock



Mineral Separation Procedures



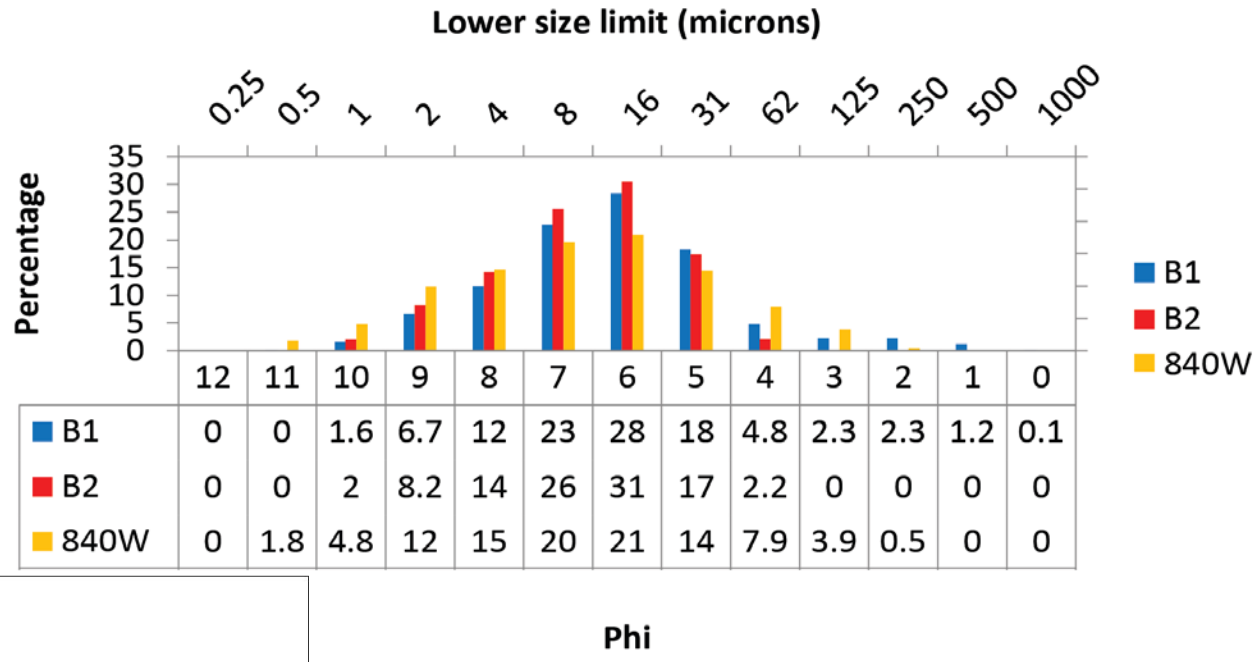
Bulk Soil Properties

Site	Sample	Munsell Color	ρ (g/cm^3)	Mean grain size (µm)	Soil Texture (USDA)
1	B1	10YR 6/3	1.5 ± 0.2	43	Silt
1	B2	10YR 5/6	1.5 ± 0.2	19.9	Silt
2	W840	7.5YR 5/4	1.8 ± 0.2	29.3	Silt Loam

Bulk Properties

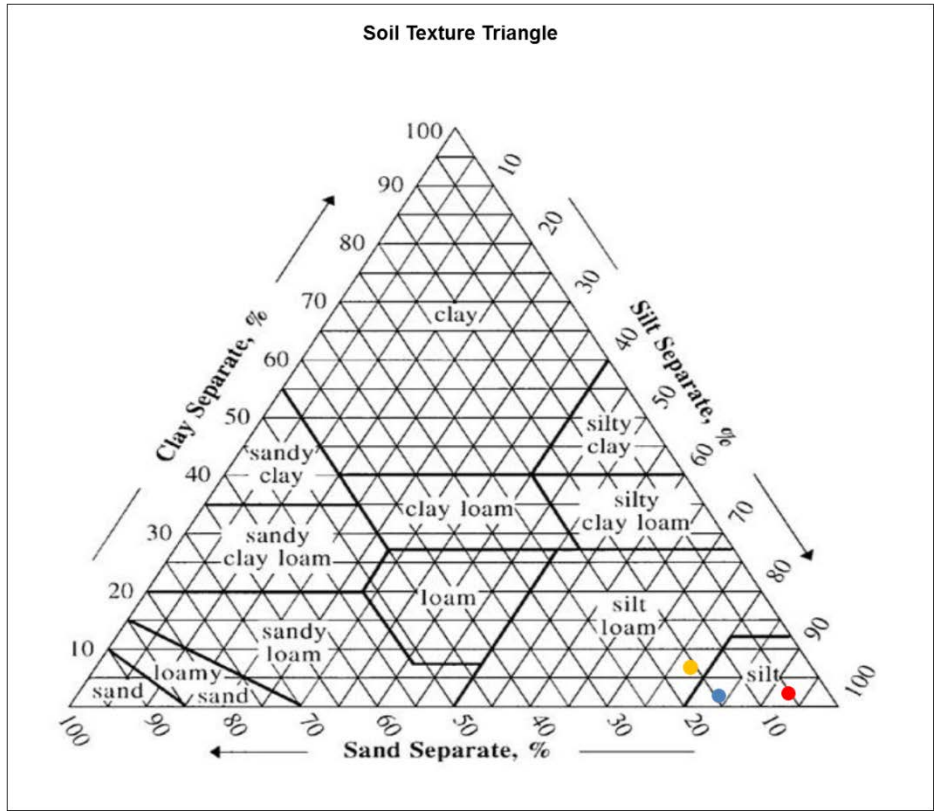
Sample	% Insol. Residue	$\epsilon_{Zr, w}$	CIA %	% OC
B1	NA	0.81	73	0.69
B2	NA	0.76	84	0.63
W840	NA	0.68	76	1.81
Mfp	99%	NA	78	
Oh	28%	NA	70	
Obc	15%		60	
PL	82%			3.92
RS	99%			0.25
HR	99%			1.02

App. DR1



	12	11	10	9	8	7	6	5	4	3	2	1	0
B1	0	0	1.6	6.7	12	23	28	18	4.8	2.3	2.3	1.2	0.1
B2	0	0	2	8.2	14	26	31	17	2.2	0	0	0	0
840W	0	1.8	4.8	12	15	20	21	14	7.9	3.9	0.5	0	0

Phi



Major element compositions

Original name	Loveland loess	Roxana silt	Peoria loess	Harpeth R.	Mfp	B1	B2	Oh	Obc	W840
SiO2	88.92	77.26	69.46	81.87	87.33	86.14	78.12	24.2	5.67	67.32
TiO2	0.32	0.87	0.71	0.99	0.33	1.03	0.99	0.2	0.04	0.64
Al2O3	5.29	10.05	9.10	7.16	5.13	5.14	8.76	4.02	1.04	6.34
TFe2O3	1.75	3.91	3.21	3.42	2.58	2.36	4.39	1.62	0.49	3.21
MnO	<0.01	0.06	0.08	0.04	<0.01	0.02	0.02	0.16	0.05	0.15
MgO	0.21	0.68	2.83	0.37	0.46	0.24	0.43	1.19	2.68	0.47
CaO	0.48	0.49	3.89	0.39	0.19	0.1	0.09	36.57	48.2	7.7
Na2O	0.04	0.48	1.33	0.32	0.09	0.37	0.27	0.13	0.11	0.25
K2O	0.19	1.66	2.07	1.1	1.05	1.03	0.95	1.19	0.3	1.09
P2O5	0.03	0.06	0.11	0.24	0.05	0.04	0.08	1.69	2.81	5.96
LOI	2.61	4.30	7.17	4.29	2.97	3.31	5.76	29.28	38.31	6.75
TOTAL	99.84	99.81	99.95	100.19	100.18	99.78	99.86	100.25	99.7	99.88

Powder XRD Results

- W840 soil has same minerals as Oh bedrock
- B1 and B2 soils do not have the same minerals as Mfp bedrock

Sample	Quartz	Microcline	Albite	HAp	Chlorite	Amphibole	Sum
B1	95	4	1	0	0	0	100
B2	95	2	1	0	2	0	100
Mfp	100	0	0	0	0	0	100
Obc	38	8	11	35	8	0	100
W840	92	2	2	4	0	0	100
Oh	87	3	1	9	0	0	100
HR	97	2	1	0	0	0	100
PL	90	2	5	0	2	1	100
RS	97	2	1	0	0	0	100
LL	98	0	0	0	2	0	100

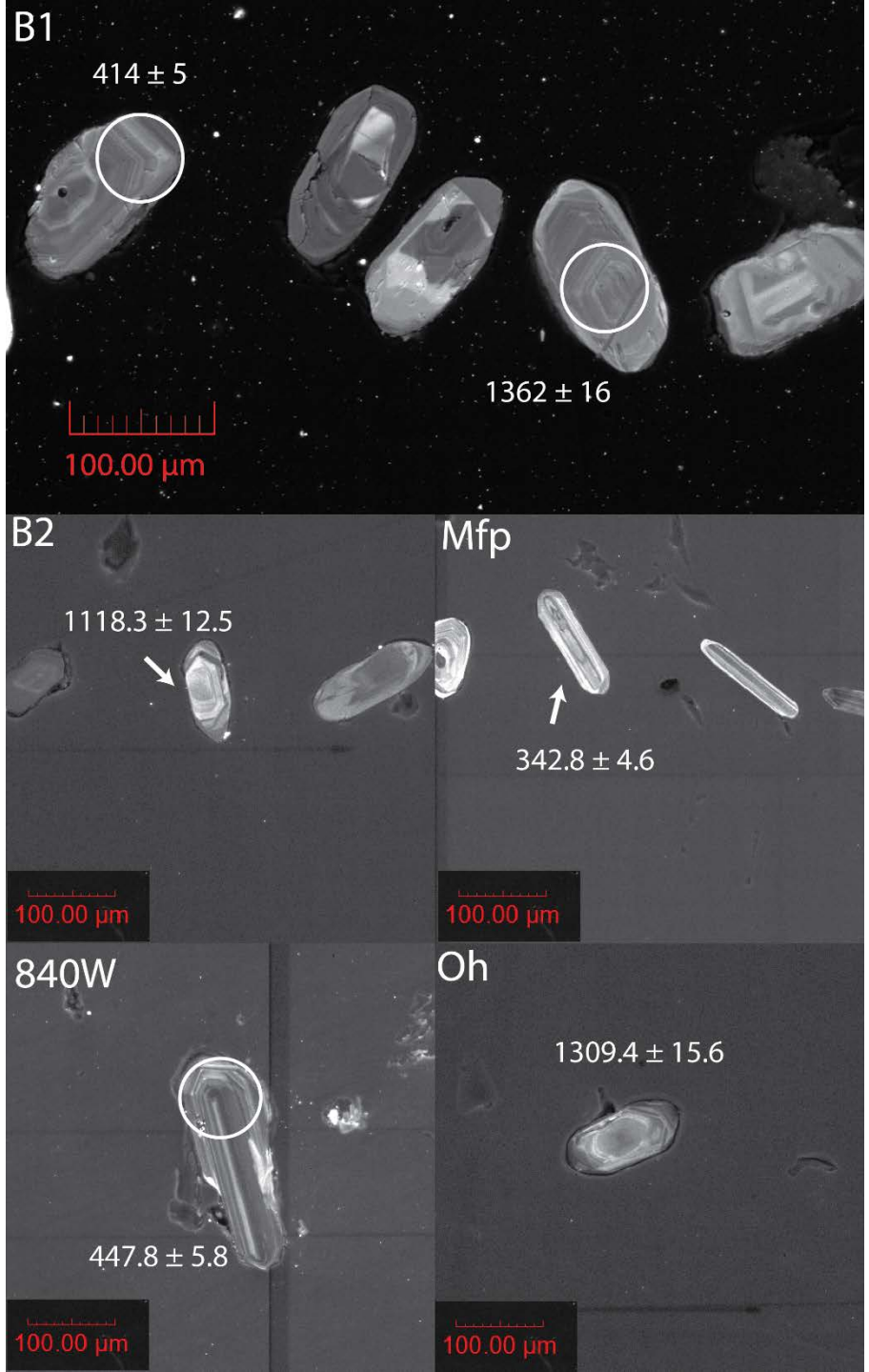
Geochronological Methods

- CL imaging – zoning and inclusions
- ThermoFisher iCAP Qc quadrupole ICP-MS with CETAC autosampler and Photon Machine Excite 193nm excimer laser ablation system
- 50 x 50 µm spot size
- Data processed in Glitter, then ET_Redux v. 3.6.25
- Analyses that were > 20% positively or negatively discordant were rejected, where % discordance was calculated as

  $$100-(100*(^{206}Pb/^{238}U \text{ date})/(^{207}Pb/^{206}Pb \text{ date}))$$

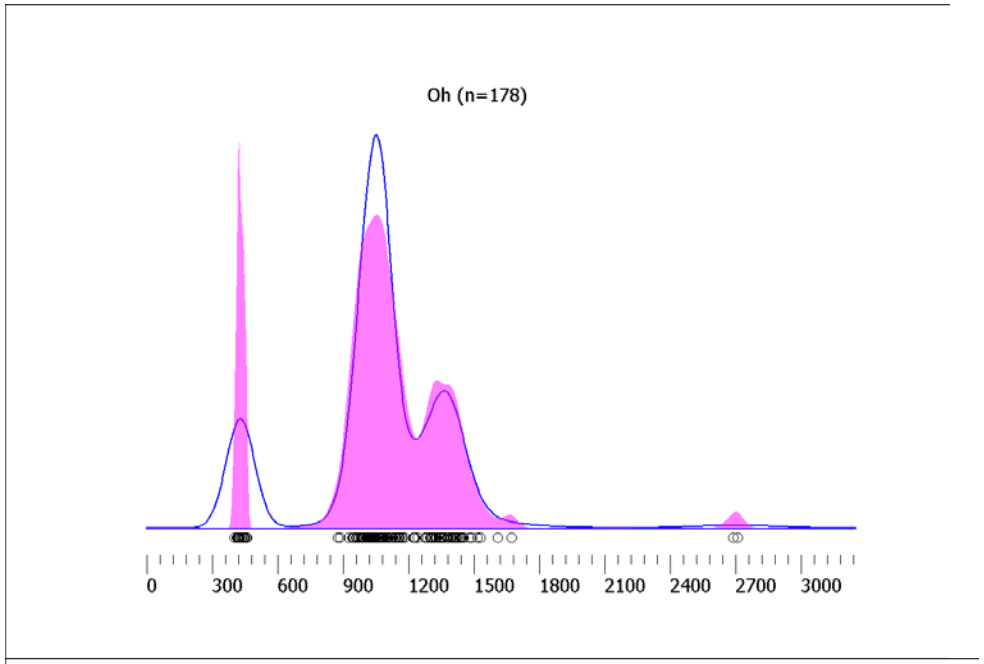
- Age spectra plotted as kernel density estimates using “DensityPlotter” program of Vermeesch (2012)

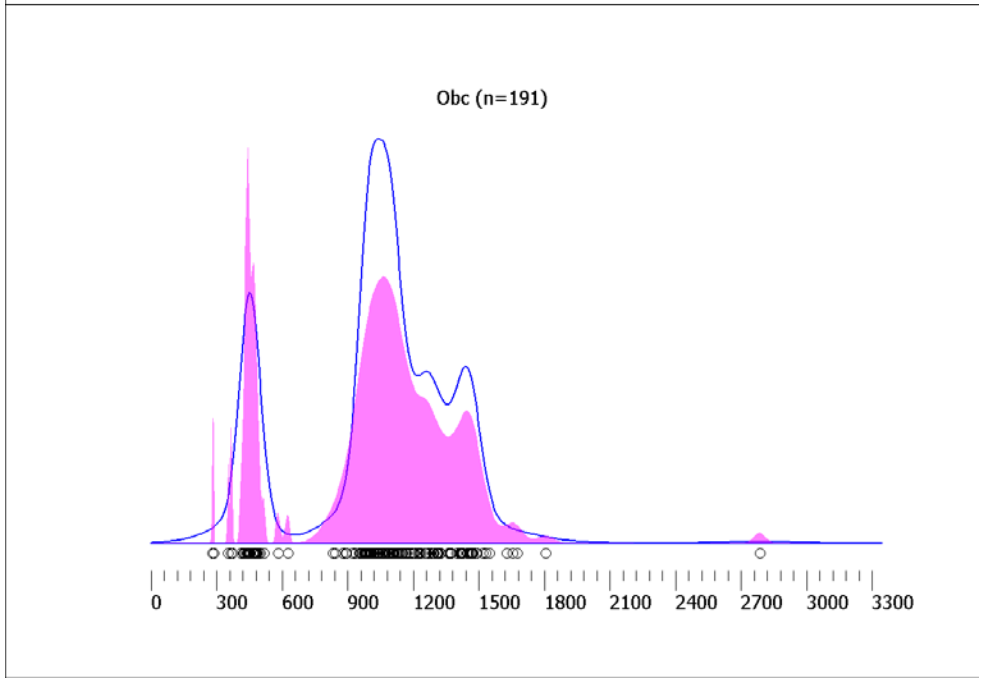
Zircon CL Images

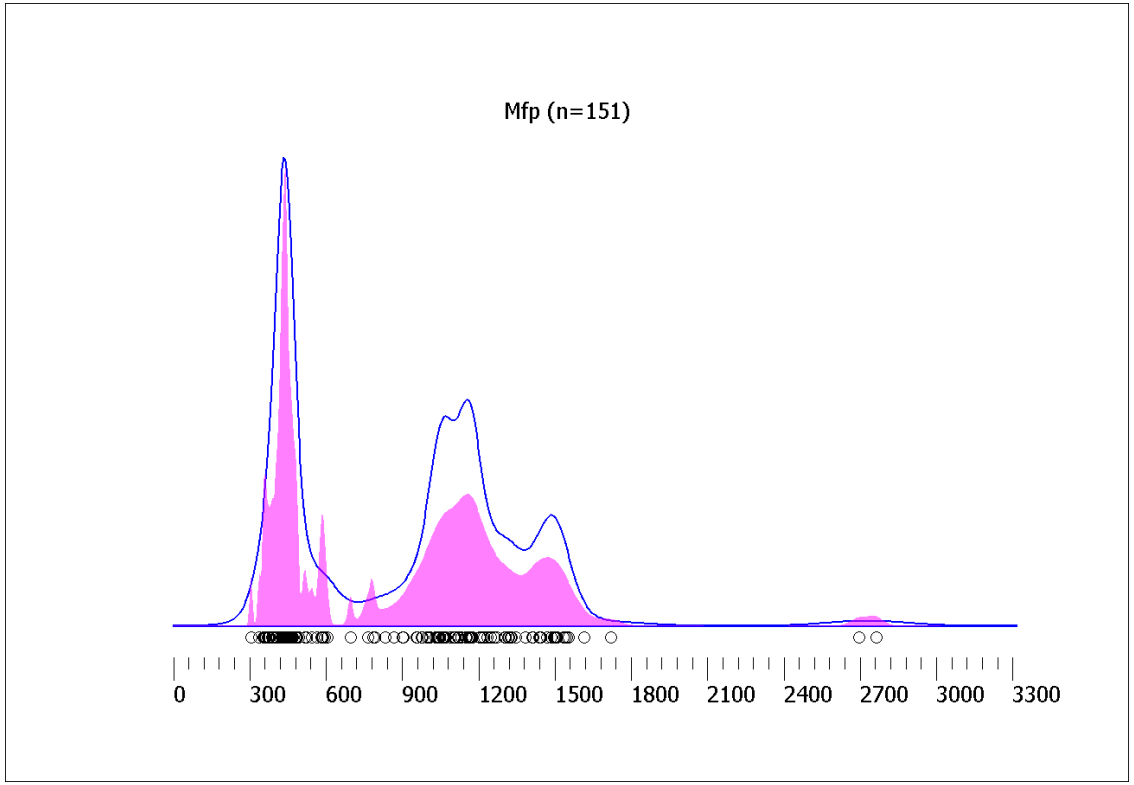


Bedrock Units

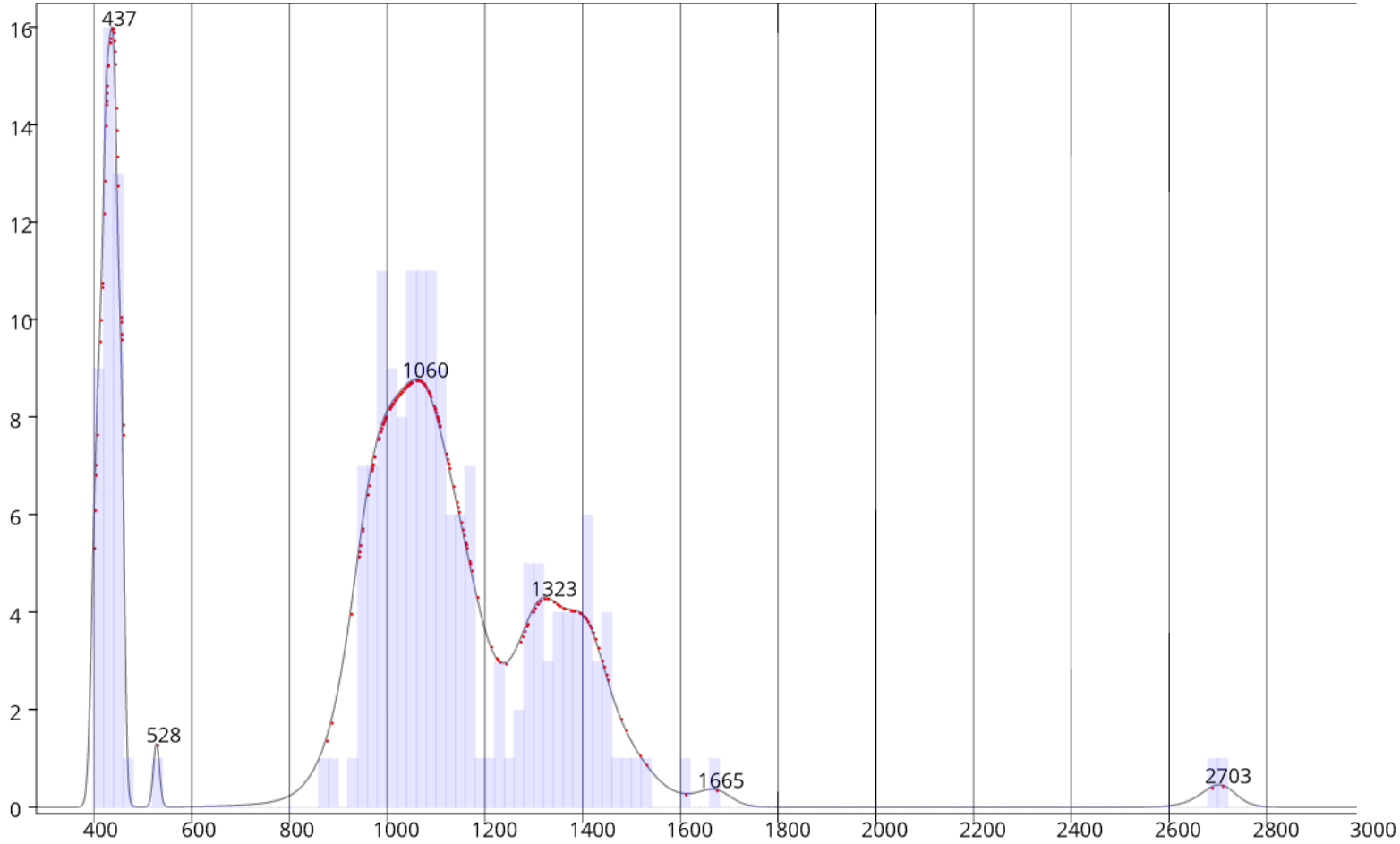
- Taconic orogeny: 430-450 Ma
- 1000-1100 Ma ages consistent with previous ages from southern Appalachians







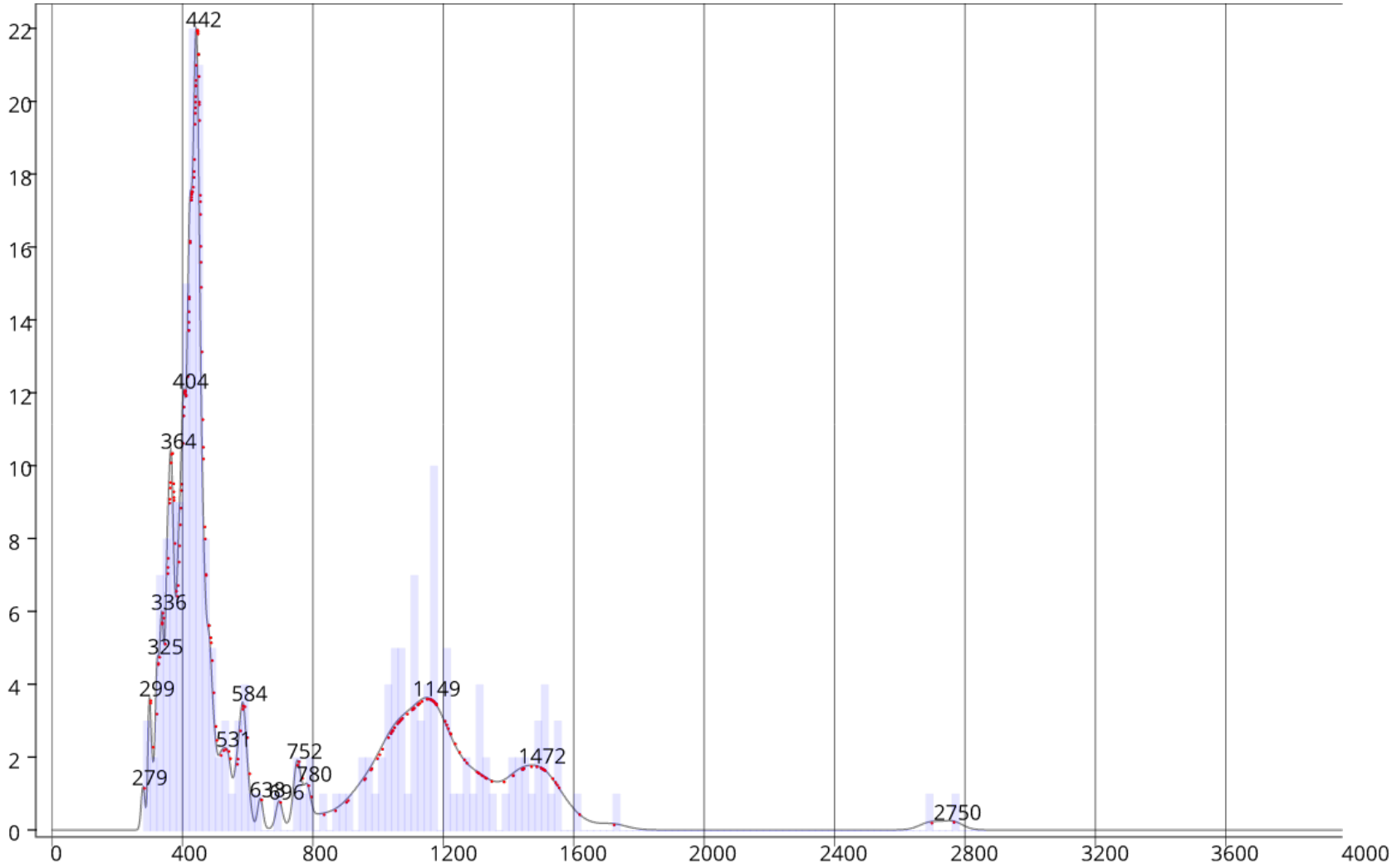
Pink represent probability density function, blue line represents the kernel density estimate (Vermeesch, 2012).



Tectonic Setting: Oh sediments

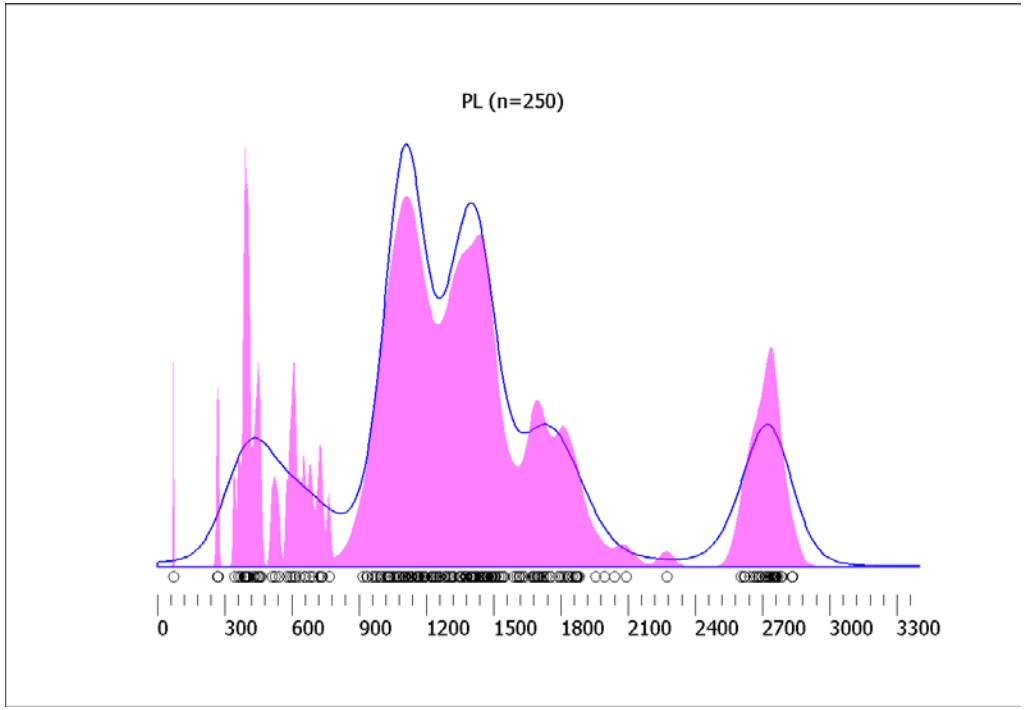
- Oh depositional age 453 Ma (Holland and Patzkowsky, 1997)
- Youngest age peak close to depositional age
- Consistent with a convergent plate margin (supra-subduction zone) setting during the Taconic orogeny.

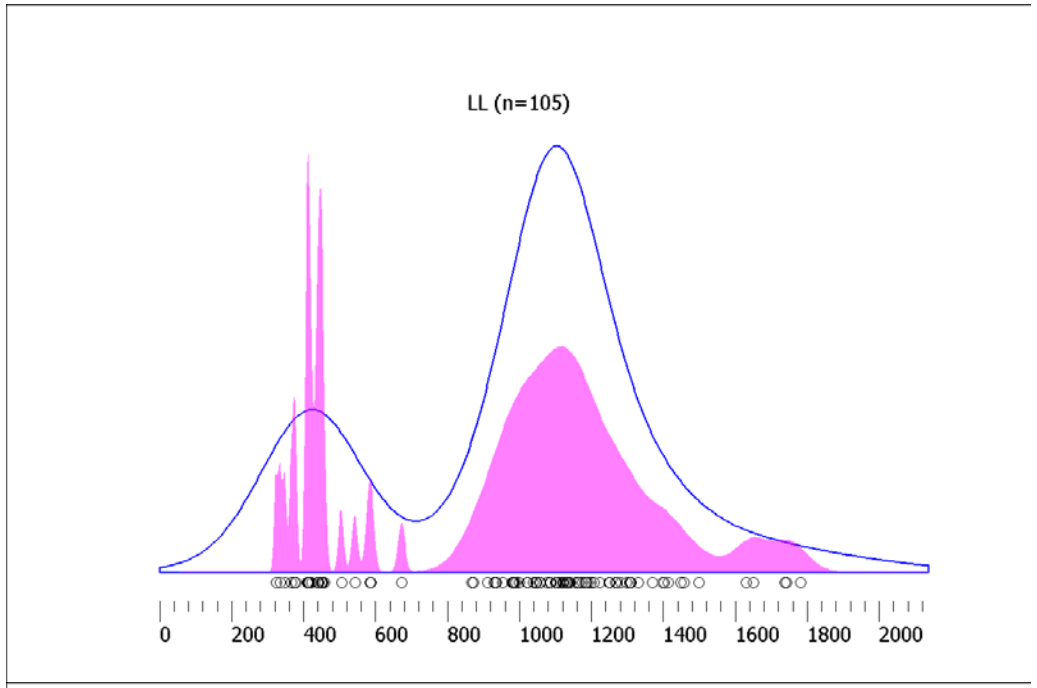
Tectonic Setting: Mfp Sediments

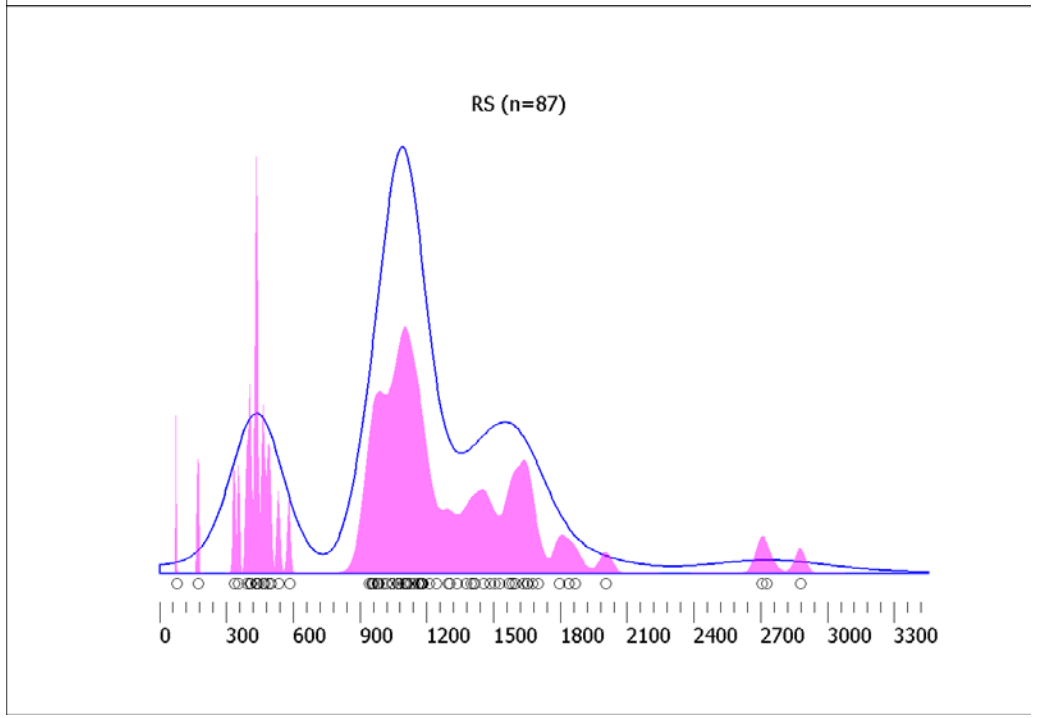


- Mfp is lower Mississippian, so depositional age ~ 350 Ma
- Youngest prominent age peak ~440 Ma
- Zircon crystals deposited ~90 Ma after crystallization
- Suggests a collisional setting, specifically the foreland basin of the Appalachians (Cawood et al., 2012).

Loess samples







Kolmogorov-Smirnov test P values

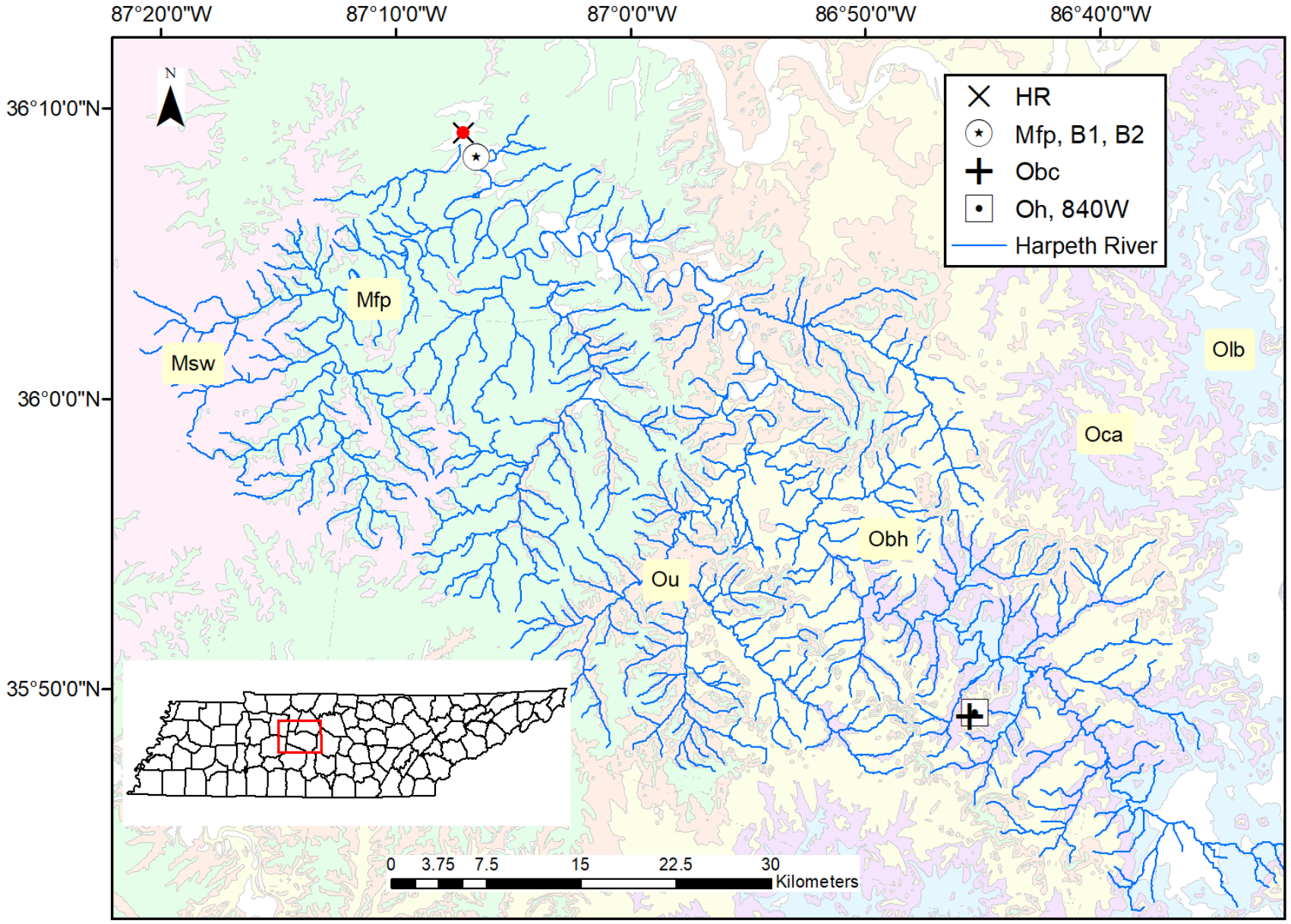
	Mfp	B1	B2	Oh	W840	LL	RS	PL	HR	Obc
Mfp	1	6.8E-07	0.001	1.0E-06	0.01	0.03	0.004	1.1E-09	8.3E-05	0.0003
B1		1	0.16	0.38	0.11	0.02	0.045	3.2E-08	0.10	0.06
B2			1	0.13	0.02	0.14	0.092	2.6E-07	0.15	0.36
Oh				1	0.37	0.09	0.013	1.4E-09	0.83	0.53
W840					1	0.49	0.034	2.1E-11	0.65	0.33
LL						1	0.073	7.1E-09	0.37	0.68
RS							1	0.0020	0.01	0.05
PL								1	1.8E-09	4.3E-09
HR									1	0.90
Obc										1

Sample Locations

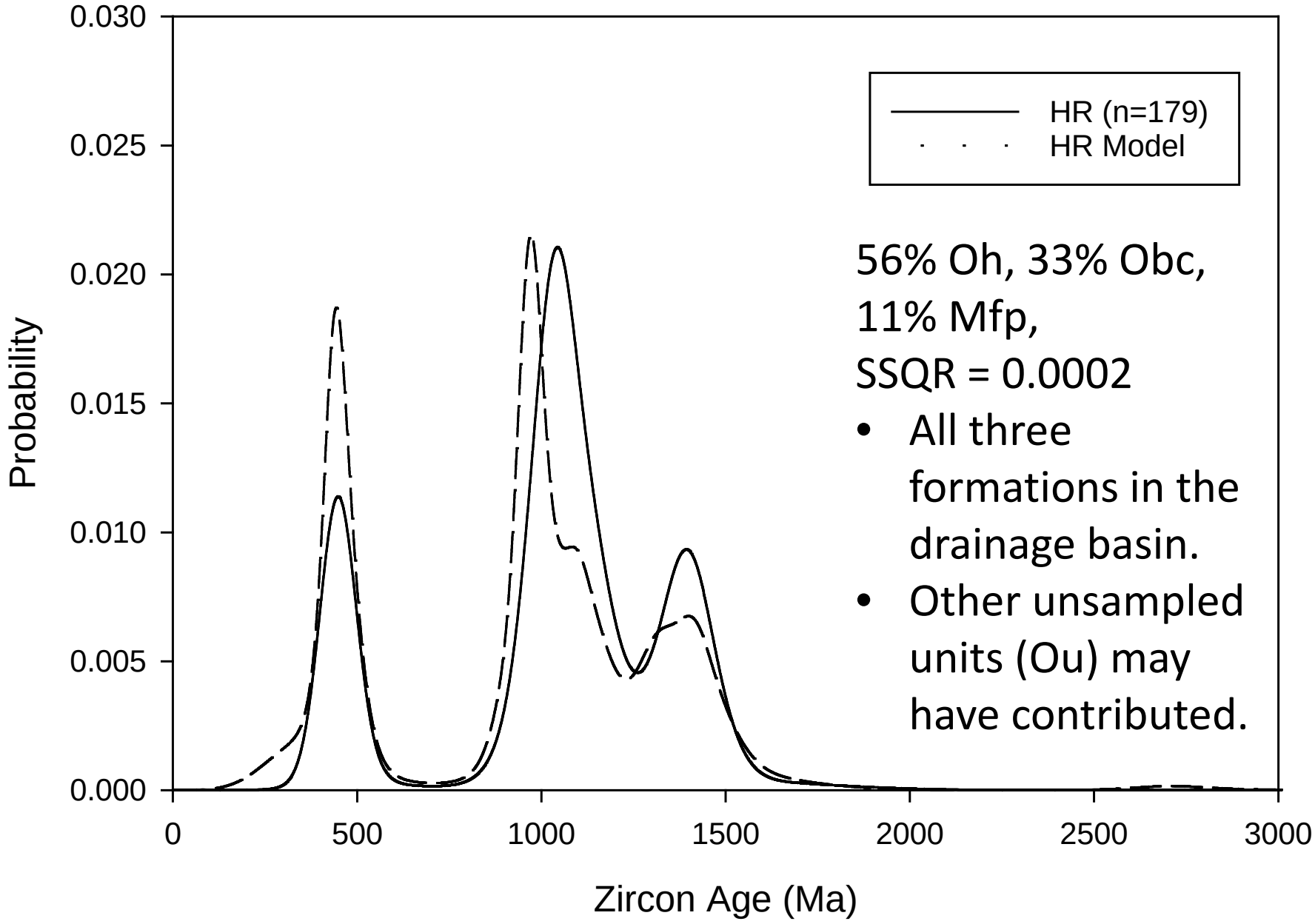
- Mfp = Mississippian Fort Payne (Site 1)
- Obh = Ordovician Bigby Cannon and

Hermitage limestones (site 2)

Site 1 located on a Q3 river terrace > 100 feet above the present Harpeth River.



Harpeth River Age Spectra

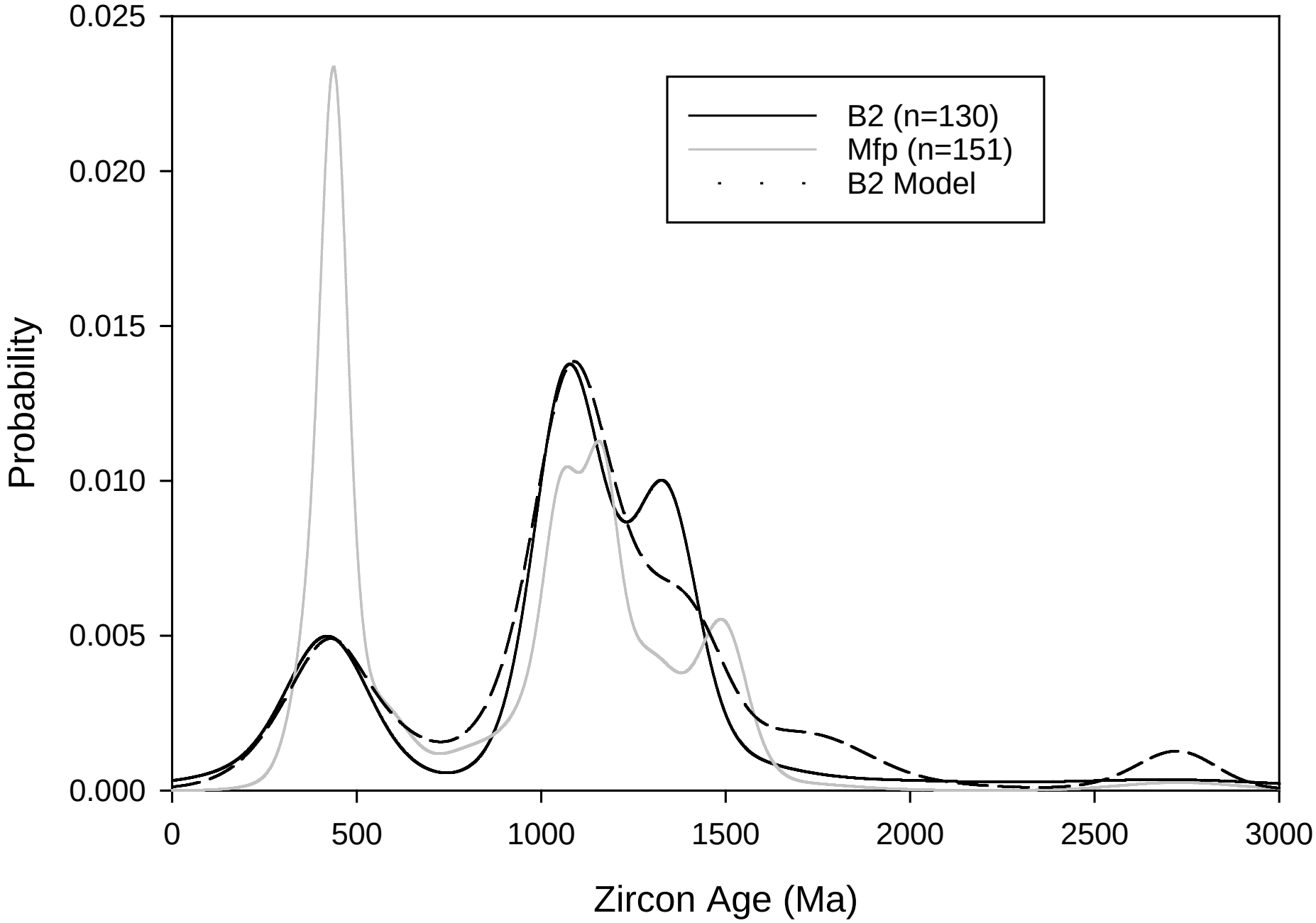


56% Oh, 33% Obc,
11% Mfp,
SSQR = 0.0002

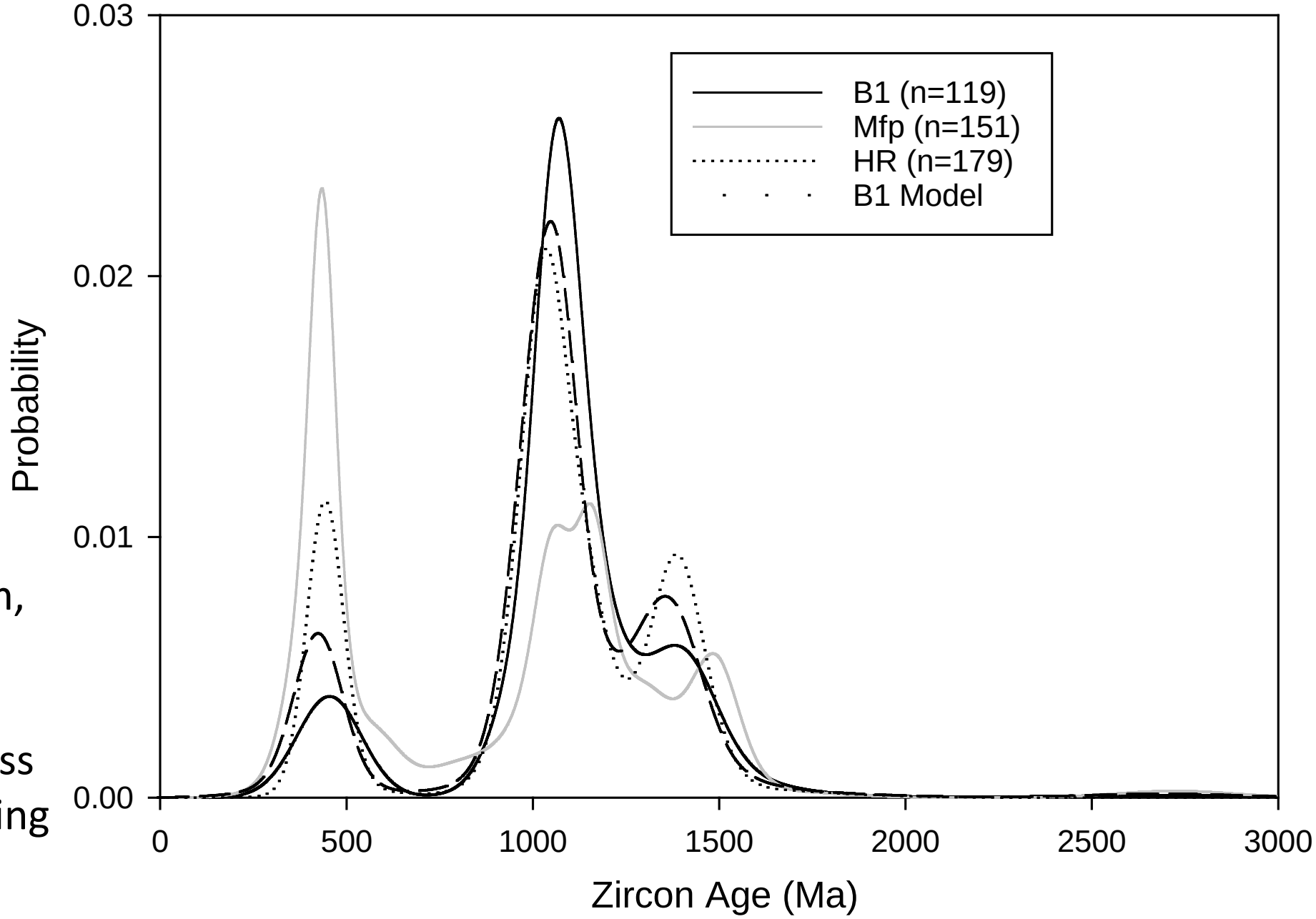
- All three formations in the drainage basin.
- Other unsampled units (Ou) may have contributed.

Deep layer B2 Age Spectra

47% LL, 29% PL, 24% Oh
SSQR = 0.00028
Since Oh is deep beneath the surface at site 1, must have been deposited as HR alluvium.

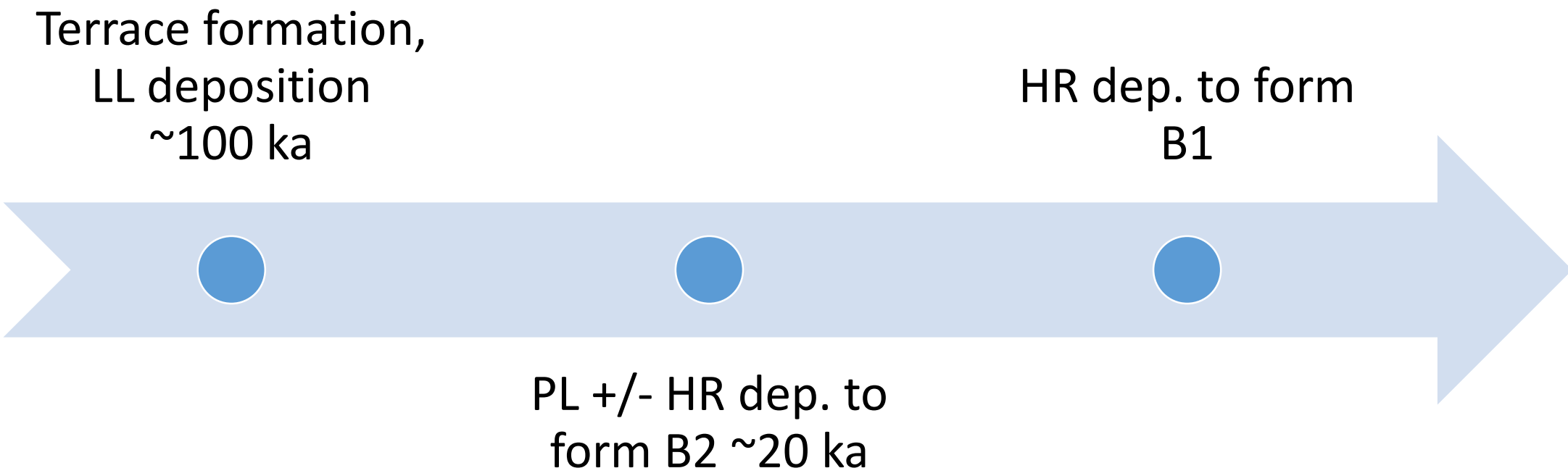


Shallow layer B1 Age Spectra



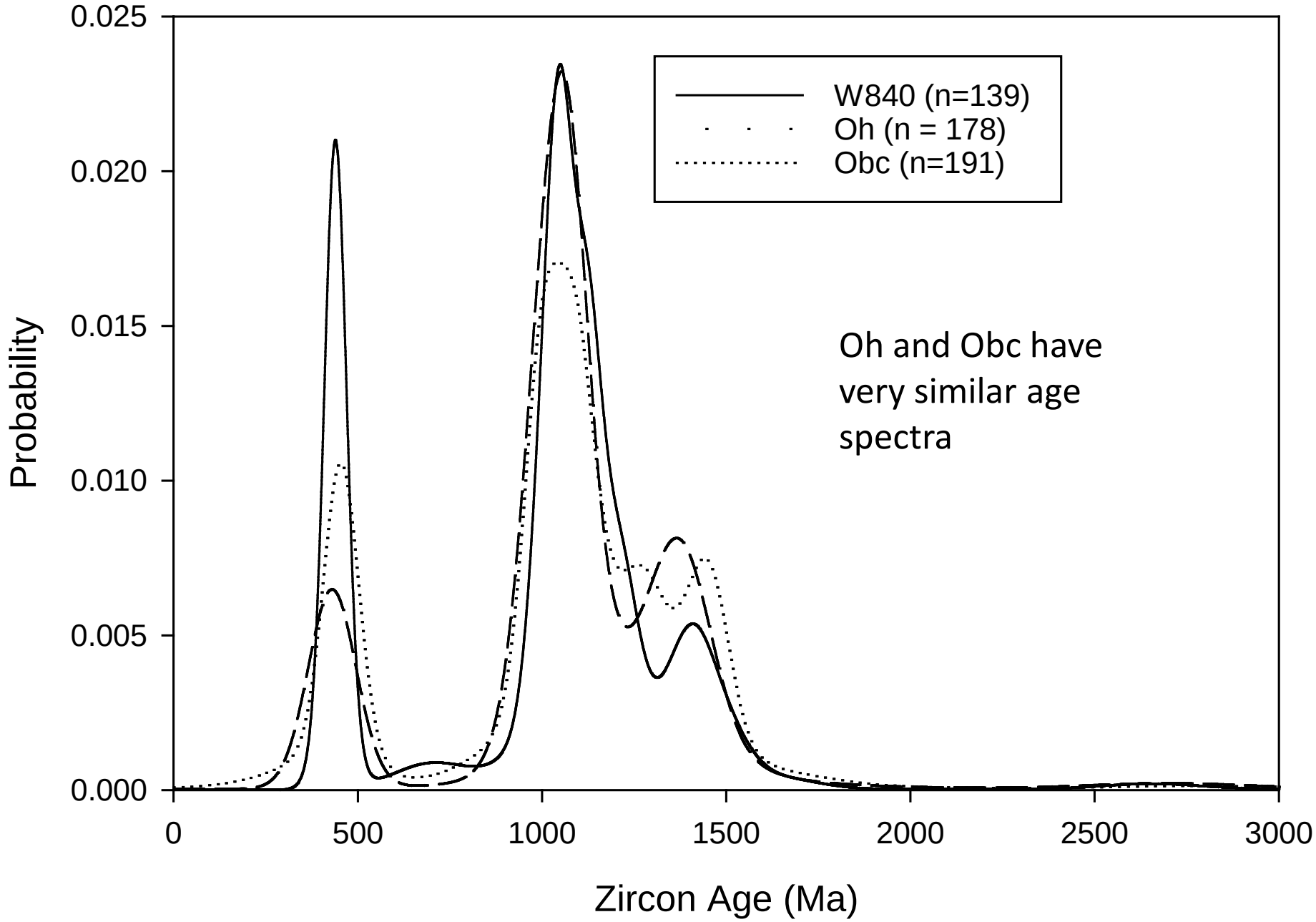
90% Oh as HR alluvium, 10% LL
SSQR = 0.0006
LL seems unlikely unless there was vertical mixing

Timeline for Site 1



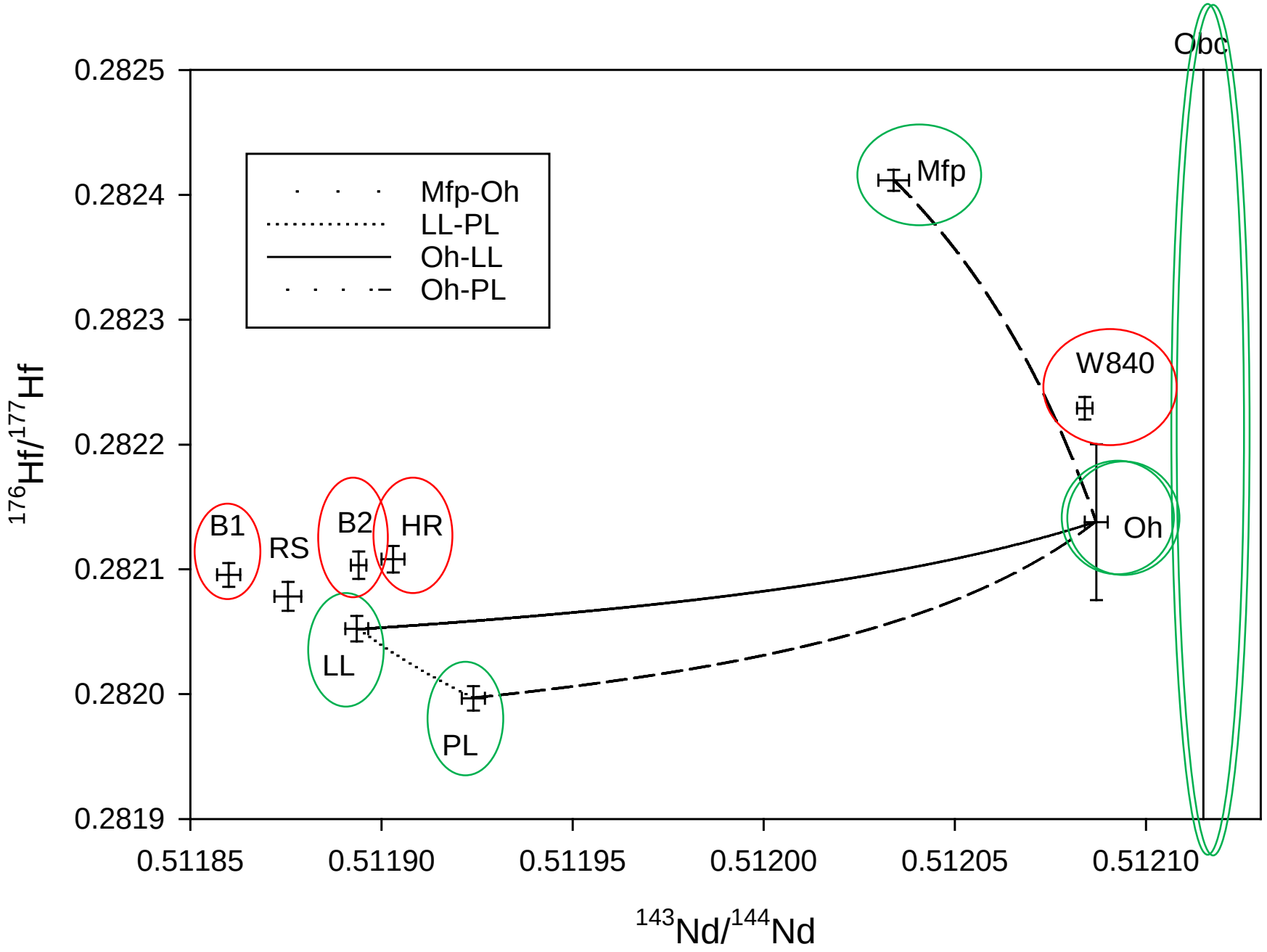
Site 2 W840 Soil Age Spectra

82% Obc,
18% Oh ± Mfp
SSQR = 0.001

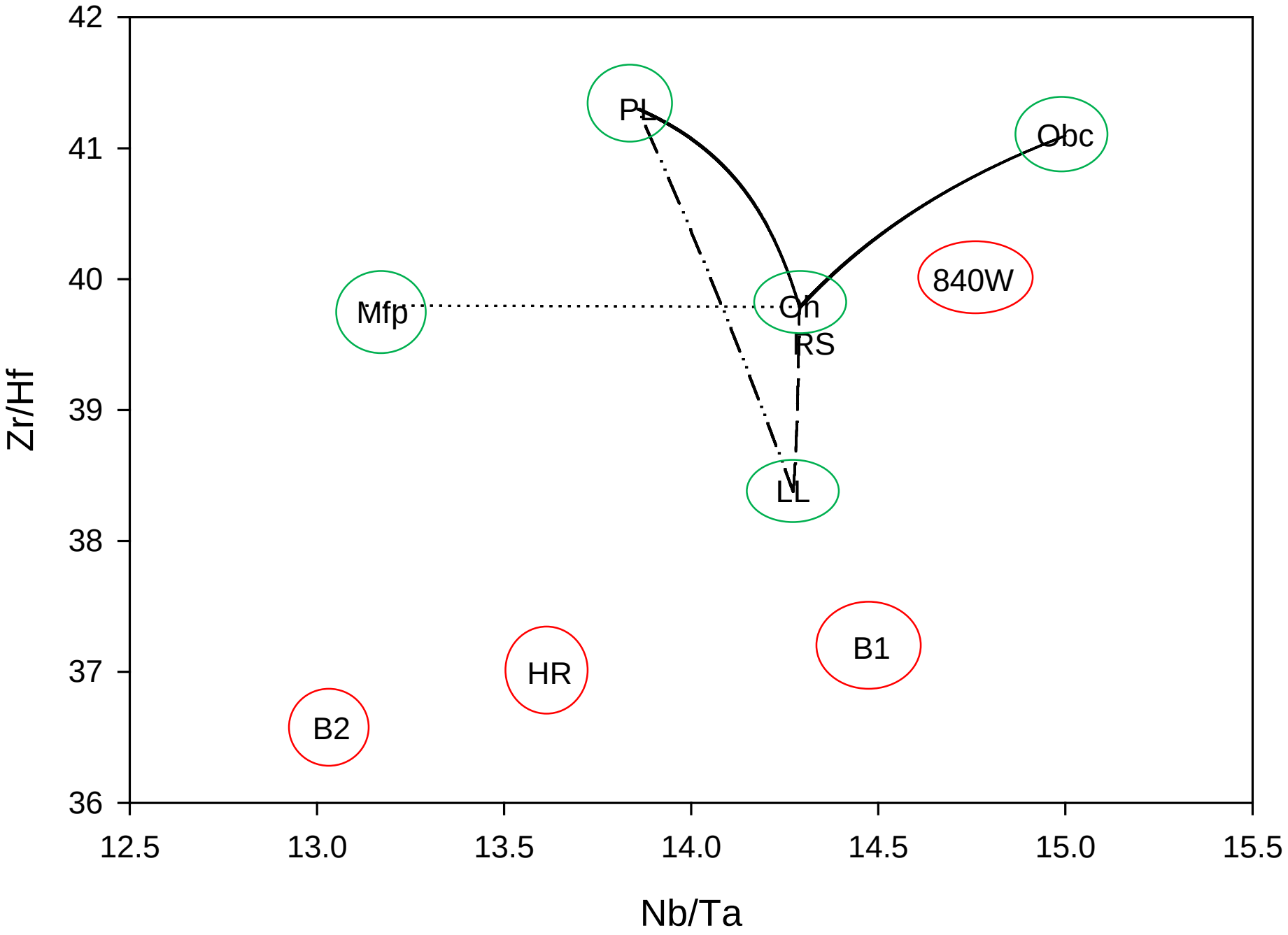


Bulk Isotope Results

B1 and B2 not close to Mfp



Trace Element Results



Evidence Consistent With Genetic Relationship?

Parent-soil pair	Mineralogy	U-Pb age spectra	$^{143}Nd/^{144}Nd$	$^{176}Hf/^{177}Hf$	Zr/ Hf	Nb/ Ta	Gd_N/Yb_N	Eu/ Eu*	τ_j §
Mfp-B1	N	N	N	N	N	N	Tie†	N	Y
Mfp-B2	N	N	N	N	N	Y	Tie	N	Y
HR-B1	Y	Y	N	Y	Y	N	Tie	Y	
HR-B2	Y	Y	Y	Y	Y	N	Tie	Y	
Oh-W840	Y	Y	Y	Y	Y	?	Y	Y	Y
Obc-W840	?	Y	N		?	?	N	N	Y
HR-W840	N	Y	N	N	N	N	N	N	
Binary Mixtures									
(Oh+LL)-B1	N	Y	N	Y	N	Y	N	N	
(Obc+Oh)-W840	?	Y	N	Y	Y	Y	Tie	Tie	
(PL+LL)-B2	Y	Y	Y	N	N	N	N	N	
(Mfp+Oh)-W840	Y	N	Y	Y	Y	N	Tie	Tie	

† "Y" if parent is closest of all parents to soil, or to binary mixing lines for pairs of parents. For mixtures, soil must plot in-between sources. "?" indicates ambiguous results. "Tie": Mfp and HR have equal GdN/YbN, and W840 has almost identical GdN/YbN and EuN/Eu*N as Oh.

§ Mass transfer coefficient τ calculated only for soil-bedrock pairs.

Conclusions

- Bedrock age peaks indicate source rocks for sediments formed during the Taconic orogeny, but many older zircons preserved from prior orogenies.
- Above Fort Payne chert bedrock, ultisol with exotic source:
  - Deposition of Loess and Harpeth River alluvium to form ultisol B2 soil horizon + intense weathering.
  - Deposition of Harpeth River alluvium to form ultisol B1 horizon.
- In-situ weathering of Oh + Obc to form W840 alfisol

Use of zircon U-Pb ages for soil provenance

- Zircon is recoverable from limestone.
- Zircon U-Pb geochronology is an effective tool for determining provenance of soil.
- However, best match in zircon U-Pb mixing model does not always agree with other datasets such as Nd and Hf isotopes.
  - Similar age spectra of our endmembers: all rocks formed from Appalachian sediments.
  - Harpeth River alluvium derived by erosion of those rocks with similar age spectra.
  - Loess deposits have many age peaks, some of which overlap with Appalachian sediment age peaks.